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北半球夏季急流遥相关: 进展与展望

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摘 要 急流遥相关是沿急流波导传播的准静止行星尺度波列,在大气环流的维持和演变中有重要作用,其异常活动常伴随极端天气气候事件的发生。研究急流遥相关在不同时间尺度的变异特征和机理不仅可以加深对大气环流变化规律的理解,还可以为相应时间尺度上的天气预报、气候预测乃至气候预估提供科学依据。根据波导的性质,北半球夏季的急流波导可以分为以正压效应为主要维持机制的副热带急流波导和以斜压效应为维持机制的极锋急流波导。本文聚焦夏季急流遥相关活动最为活跃的欧亚大陆地区,在回顾急流遥相关理论的基础上,简要总结了近年来关于夏季副热带急流遥相关和极锋急流遥相关的形成条件、动力机制、气候影响和未来变化的研究进展,并展望了一些未来值得继续深入研究的问题。

关键词: 急流; 大气波导; 大气遥相关; 行星波; 波流相互作用
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Abstract Waveguide teleconnections are quasi-stationary planetary waves propagating along the jet stream waveguide. They play an essential role in the maintenance and evolution of atmospheric circulation, and their abnormal activities are often associated with extreme weather and climate events. Studying the variation and mechanism of waveguide teleconnections at different timescales can not only deepen our understanding of the dynamics of atmospheric circulation but also serve as a scientific basis for weather forecasting, climate prediction, and even climate projection at corresponding timescales. According to the dynamical property of the jet stream waveguide, the atmospheric waveguide can be divided into the subtropical jet waveguide, which is maintained essentially by the barotropic effect, and the polar front jet waveguide, which is primarily maintained by the baroclinic effect. This paper focuses on the waveguide teleconnections in the boreal summer over the Eurasian region, when and where waveguide teleconnections are the most active. It first reviews the history and theory of waveguide teleconnection and then summarizes some recent advances in waveguide teleconnections along the subtropical jet and polar front jet, including their formation conditions, dynamic mechanisms, climate impacts, and future changes. Finally, it discusses some possible issues that deserve future investigation.

Keywords Jet stream, Atmospheric waveguide, Atmospheric teleconnection, Planetary wave, Wave-mean flow interaction

1 引言

急流遥相关也称为波导型遥相关(waveguide teleconnection),是沿急流波导传播的准静止大气 Rossby

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作者简介 王林, 男, 1981 年出生, 研究员, 主要从事气候动力学研究。E-mail: wanglin@mail.iap.ac.cn **资助项目** 国家自然科学基金项目 41925020、42005057、42261134532 **Funded by** National Natural Science Foundation of China (Grants 41925020, 42005057, 42261134532) 波(Ambaum 1997; Branstator 2002; Ding and Wang 2005; White et al. 2022)。与太平洋-北美遥相关和北大 西洋涛动等具有显著经向分量的大气遥相关不同,急流遥相关主要沿西风急流在纬圈方向传播,因此经向 尺度小、纬向尺度大,有时甚至能环地球一周(谭本馗和陈文 2014; Branstator and Teng 2017)。从全球尺 度看,急流遥相关能引发大至全球尺度的急流异常,造成大气、海洋和海冰等地球气候系统成员在高、低 纬度之间的动量、热量和能量交换,导致大气环流的大规模调整(布和朝鲁等 2008; 张耀存等 2008; 任 雪娟等 2010; Hoskins and Woollings 2015; Wirth et al. 2018); 从区域尺度看,急流遥相关能引发一系列沿 纬圈方向排列的异常高压和低压系统的有组织增长或衰亡,从而深刻影响中高纬度不同地区的天气和气候 (Schubert et al. 2011; Trenberth and Fasullo 2012; Palmer et al. 2013; Wolf et al. 2018), 特别是导致相距上 千公里的不同地区同时出现极端天气气候事件(Screen and Simmonds 2014; 叶丹和张耀存 2014; Huang et al. 2015, 2017; Rodrigues and Woollings 2017; Lin and Yuan 2022)。相较于单个发生的极端事件,急流遥相关 在不同地区同时引发的多个极端事件往往会造成更为严重和复杂的社会影响(Zscheischler et al. 2018; Raymond et al. 2020)。例如,当多个农业区同时遭受极端天气气候事件时会严重威胁全球的粮食产量,甚至 引发全球范围的粮食危机(Tigchelaar et al. 2018; Gaupp et al. 2020; Kornhuber et al. 2020; Raymond et al. 2020; Rogers et al. 2022);而多个同时发生的极端天气气候事件还会造成再保险行业的赔付增加,并将潜在的金融 风险传播至全球的不同地区和行业(Swiss Re 2019a, 2019b; Raymond et al. 2020)。因此,近年来关于急流 遥相关特征和影响的研究受到社会各界的关注。

夏季,随着极地和低纬度地区的温度差异变小,中高纬度大气盛行的西风减速,急流分化为极锋急流和副热带急流两支(Lee and Kim 2003; Kuang and Zhang 2005; Huang et al. 2014)。这种独特的双急流结构使得急流遥相关的活动在夏季达到全年最强,对中高纬大气环流的影响也最大(Lin 2014; Wang et al. 2014; McKinnon et al. 2016; Coumou et al. 2018; Rousi et al. 2022)。例如,2010年夏季,副热带急流和极锋急流上的大气遥相关均异常活跃,在欧亚大陆上空表现为南北两支的纬向双波列结构(图 1a)。与此相伴随,俄罗斯西部、蒙古、我国华北和东西伯利亚地区几乎同时遭受到极端高温事件的侵袭(图 1d),其中俄罗斯西部的极端高温最为严重,导致约 55000 人死亡,粮食产量减少近 30%,直接经济损失 150 亿美元;同时巴基斯坦遭遇严重的洪水灾害,导致约 2000 万人受灾,直接经济损失 430 亿美元(Barriopedro et al. 2011; Trenberth and Fasullo 2012; Teng and Branstator 2019; Xu et al. 2021a)。2019年和 2020年夏季,相似的大气环流型再次出现(图 1b-1c),欧亚大陆多地频繁遭受极端天气事件的侵袭(图 1e-1f),造成严重的人员伤亡和社会财产损失(Sousa et al. 2020)。可见,深入揭示夏季急流遥相关的变异机理及其气候影响,对认识中高纬大气环流的变化规律和极端天气气候事件的预报预测均有重要科学意义和潜在应用价值。

自从急流波导的概念提出以来,已有很多学者研究了副热带急流上遥相关的变异机理和气候影响。近 年来,随着极锋急流的波导作用得到认识,国内外学者对极锋急流遥相关进行了深入的研究,取得了许多 新的研究成果。本文将聚焦北半球夏季和欧亚大陆地区,简要回顾近年来关于副热带和极锋急流遥相关变 异机理和气候影响方面的进展。

2 急流波导理论的发展

大气波导的概念最早可以追溯至 20 世纪 60 年代。Charney and Drazin(1961)和 Eliassen and Palm(1961) 在研究行星波沿垂直方向的能量传播时从理论上指出,准定常行星波只能在有垂直切变的背景西风中传播。 Dickinson (1968)和 Matusno (1970)发现准定常行星波在从对流层上传至平流层过程中总倾向于向平流层 的极夜急流传播,并提出了"行星波波导"的概念,以描述行星波在垂直上传的过程中倾向于向急流附近 的风速大值区传播的特征。随后,气象学家们将波导的概念运用到行星波在β平面或二维球面的传播问题 中(Longuest-Higgins 1964, 1965; Hoskins et al.1977; Hoskins and Karoly 1981),特别是 Branstator (1983) 基于理想正压模式讨论了准定常行星波在急流附近的水平传播特征,发现急流附近的准定常行星波能量并 非沿着 Hoskins and Karoly (1981)提出的大圆路径传播,而是沿着急流在纬圈方向传播,从而首次在模式 中给出了急流可以作为 Rossby 波波导的启示。20 世纪 80 年代是大气遥相关研究蓬勃发展的时代,但以 Wallace and Gutzler(1981)为代表的学者们一直没有在观测资料中发现急流波导的事实证据,这可能与当时的研究主要使用对流层中层的位势高度场从而无法抓住空间尺度更小的急流遥相关有关。直到 20 世纪 90 年代,Hsu and Lin(1992)首次从 250hPa 流函数场中发现一个位于副热带急流上的源自北美、延伸至欧亚大陆南部的定常波列,从而给出了急流遥相关存在的观测证据。随后,Hoskins and Ambrizzi(1993)从理论上解释了准定常行星波沿急流传播的原理,从而奠定了急流遥相关的理论基础。

与光学中的 Snell 定律类似, Rossby 波作为一种在地球流体介质中传播的波动会朝着折射指数大的区域 传播 (Platzman 1968; Rhines 1979; Held 1983; 杨菘等 2015; Y. Li et al. 2021)。在准地转近似下, Rossby 波的折射指数可表示为 $K_s^2 = \frac{\bar{q}_y}{\bar{u}_a} - \frac{\varepsilon}{4H^2}$ (Andrews et al. 1987),其中 $\varepsilon = f_0^2/N^2$, q为准地转位涡,H为大气标 高, \bar{u}_a 为地转纬向风, f_0 为科氏参数,N为浮力频率。Hoskins and Ambrizzi(1993)将该问题简化为正压大 气,得到折射指数 K_s^2 正比于绝对涡度的经向梯度 $\beta_* = \beta - \frac{\partial^2 \overline{u}}{\partial v^2}$,其中 β 为科氏参数的经向变化。 β_* 主要依赖 于纬向风的经向结构,并在副热带急流轴附近有极大值,这表明副热带急流是 Rossby 波传播的有效波导。 因此,特定波数的行星波在副热带急流轴附近传播时会不断在急流波导内发生反射,从而使波动能量只能 沿着副热带急流在东西方向传播(Chang and Yu 1999; Lee and Feldstein 1996; Martius et al. 2010),此时行 星波可近似看作是以群速度 $C_g = \bar{u} + \frac{\beta_*}{k^2}$ 在一维系统中传播。一方面,这意味着沿急流传播的准定常行星波 能量会快于基本气流的速度;另一方面,由于受到急流波导的约束,准定常行星波的经向能量频散被限制, 因此行星波可以在东西方向传播至非常远的区域,在冬季甚至可以环地球一周(Manola et al. 2013)。尽管 以上急流波导理论是在线性框架下得到的,且推导过程中用到的 WKB 假设在观测中并不严格成立,但其所 得的结论仍有很大的理论指导意义,一个例证就是观测中副热带急流对 Rossby 波的捕获作用。如果进一步 考虑非线性效应,上述急流波导理论仍然成立,只是行星波的波峰和波谷位置会发生有限的经向位移,且 波动在增长至一定的振幅后会发生破碎(Naoe and Matsuda 1998, 2002),但这并不会改变急流波导对行星 波经向能量频散的限制作用,从而证实了急流波导在非线性系统中的适用性(Drischel and McIntyre 2008; Methven and Barrisford 2015).

Ambrizzi et al. (1995) 在 Hoskins and Ambrizzi (1993) 之后进一步研究了北半球夏季的情况,发现夏季的急流依然存在波导效应,只是波导的强度相对冬季有所减弱且波导的位置有向高纬度的移动。从对流 层高层的β_{*}分布(图 2a)可以看到,夏季副热带急流附近有β_{*}的极大值,这与冬季的情形类似,表明副热 带急流在夏季也是一支有效的大气波导,并可以解释观测中副热带急流上的遥相关活动(图 1a-1c)。然而, 夏季极锋急流附近并不存在β_{*}的大值区(图 2a),但观测中却有活跃的大气扰动且常常可看到类似急流遥相 关的纬向波状结构(图 1a-1c),表明 Hoskins and Ambrizzi (1993)和 Ambrizzi et al. (1995)提出的急流波 导理论仍不完善。Iwao and Takahashi (2008)分析了上述理论与观测不匹配的成因,指出这是由于前人所 使用的正压大气假定忽略了大气层结对急流波导效应的贡献。Xu et al. (2022a)进一步讨论了夏季极锋急流 和副热带急流的波导作用,给出了位涡经向梯度*q_y*所表征的波导结构(图 2b),并从定量的角度指出夏季欧 亚大陆地区的极锋急流是一支独立于副热带急流的有效波导。

3 副热带急流上的大气遥相关

关于夏季副热带急流遥的研究是伴随着亚洲夏季风的研究而发展的(Wu 2017; Chowdary et al. 2019; Wang et al. 2021)。从 20 世纪 80 年代开始,人们就注意到印度夏季风和东亚夏季风之间存在显著的同位相 联系(Tao and Chen 1987; 郭其蕴和王继琴 1988),并对其成因进行了深入的分析(Kripalani and Singh 1993; Kripalani and Krlkarni 2001; Zhang 2001; Kim et al. 2002)。Krishnan and Sugi (2001)发现,日本梅雨降水 偏多的年份中,欧亚大陆中纬度地区的对流层中高层存在一支东西走向的异常波列,其活动中心分别位于 里海、蒙古和日本附近。Lu et al. (2002)和 Wu (2002)从波动本身出发,首次识别和定义了沿副热带急

流传播的波列结构,指出它是大气内部稳定存在的固有模态以及它在联系印度夏季降水和东亚夏季降水中的作用。Enomoto et al. (2003)系统分析了该波列的激发、传播和气候影响,并将其命名为"丝绸之路"遥相关 (Silk Road pattern, SRP)。SRP 与后续研究提出的"环球遥相关"(Ding and Wang 2005)在欧亚大陆上空的结构高度相似,许多研究将二者看作同一种遥相关,但也有研究注意到二者的差异并对其成因进行了讨论 (Chowdary et al. 2019; Zhou et al. 2019)。在本文中,我们不去深入探讨二者的关系,而是将 SRP 看作环球遥相关在欧亚大陆的区域性体现。

SRP 可定义为亚洲副热带急流区域(图 2a 中的红框) 200hPa 经向风年际变化的经验正交函数(EOF) 第一模态(Kosaka et al. 2009)。它在水平方向表现为沿副热带急流传播的准静止 Rossby 波,主要活动中心 分别位于欧洲、里海、蒙古和朝鲜半岛附近(图 3a-b),垂直方向呈准相当正压结构,最大振幅出现在对流 层顶附近(图 3c-d)。在静力平衡的约束下,其异常高、低压中心分别对应对流层内的暖、冷温度异常,且 温度场略微落后于高度场。相较于高度场,SRP 在涡度场上呈现更加清晰的波列结构(图 3a-b),这是因为 涡度场比高度场更能够刻画 SRP 等波数较大、波长较短的急流遥相关的空间结构。当 SRP 处于正位相时(图 3a-b),乌拉尔山、东北亚和印度地区的夏季降水偏多、气温偏低,东欧、中西伯利亚、韩国和日本的夏季 降水偏少、气温偏高(图 3e, 3f)。上述气温和降水异常与亚洲夏季风系统有紧密联系(Enomoto 2004; Kosaka et al. 2012; Hong and Lu 2016; Wang et al. 2021; Muetzelfeldt et al. 2023),反映了 SRP 对印度季风降水(Ding and Wang 2005)、长江流域降水(Hsu and Lin 2007; Wei et al. 2014, 2015)和韩国降水(Choi et al. 2016) 的影响,以及在欧亚大陆夏季极端天气事件中的作用(Li et al. 2017; Xu et al. 2022b; Na and Lu 2023)。

阐明 SRP 的维持机制对理解 SRP 的空间结构有重要意义。早期研究认为, SRP 所呈现的波列结构是某一大气外强迫源所激发的准定常行星波沿副热带急流传播的结果(Lu et al. 2002; Enomoto et al. 2003)。后续研究表明,单纯的准定常行星波传播并不足以解释 SRP 的维持,而是还需要 SRP 从背景气流中汲取不稳定能量特别是有效位能(Sato and Takahashi 2006; Kosaka et al. 2009),这说明 SRP 是行星波能量频散和大气不稳定性相结合的产物,其中斜压过程在 SRP 维持中起到关键作用。Chen et al. (2013)利用一个简化的线性三层斜压模式对此进行了深入分析,指出与 Held et al.(1986)讨论的 Rossby 外波的斜压不稳定类似,天气尺度瞬变波所引起的低层热力耗散可激发 Rossby 外波的斜压不稳定,这是 SRP 获取斜压能量的主要方式。

与维持机制不同,SRP的激发机制需要解释产生SRP的扰源。基于Rodwell and Hoskins(1996)提出的"季风-沙漠"机制,研究认为SRP的激发受到印度季风降水的控制(Enomoto et al. 2003),并与印度季风降水之间存在双向的耦合过程(Ding and Wang 2005; Ding et al. 2007; Wang et al. 2021): 一方面,印度季风降水释放的凝结潜热可通过"季风-沙漠"机制激发SRP;另一方面,当SRP的扰动能量环球一周并再次传播至印度附近时,会通过加强印度上空的环流异常从而进一步增强印度降水的异常(Wang and Xie 1996; Xie and Wang 1996)。后一过程并非只有简单的东传Rossby波起作用,而是存在季风降水激发的东传Rossby 波和西传Rossby 波的共同作用(Beverley et al. 2011)。由于印度季风降水对SRP的激发有重要作用(Beverley et al. 2019),厄尔尼诺-南方涛动(ENSO)又是影响印度季风的重要因子,因此SRP也会受到ENSO的影响(Ding et al. 2011)。此外,欧洲降水(Saeed et al. 2014;Lin et al. 2017)、印度洋降水(Chen and Huang 2012)、东亚夏季风降水(Zhou et al. 2020)、北美土壤湿度(Teng et al. 2019)以及北美、赤道非洲、特别是地中海以东的非绝热加热(Yasui and Watanabe 2010)等也对SRP的激发有一定作用。这些大气外强迫因子激发 SRP的物理过程具有一定的相似性,都是通过非绝热加热直接扰动副热带急流或是通过辐散风将激发的Rossby

尽管很多研究认为大气外强迫在 SRP 的激发中发挥了重要作用,但一些基于长时间模式积分的研究却 发现 SRP 与印度季风降水等大气外强迫之间并没有显著的联系(Teng et al. 2013),且多种大气外强迫因子 都能激发 SRP 的结果在一定程度上暗示了这些外强迫因子以外的某种因素可能在激发 SRP 过程中起到了更 关键的作用,这种因素就是大气内部的不稳定或大气与地形的相互作用等大气内动力过程。例如,以随机 的白噪声强迫正压大气模式便能形成与观测中 SRP 非常类似的主导模态(Ding et al. 2011),大气内部的自

由波可以和受到稳定地形或热力强迫激发的定常波之间发生共振从而形成 SRP (Petoukhov et al. 2013),北 大西洋涛动向下游沿副热带急流的传播也可以激发出 SRP (Hong et al. 2018)。因此,SRP 本质上是沿副热 带急流传播的大气本征模态,而印度季风降水等大气外强迫可能更多地是对这一本征模态起到调控作用。

以上关于 SRP 的研究均聚焦年际或次季节时间尺度, 近年来人们开始关注 SRP 在更长时间尺度上的变化。Wang et al. (2017)基于多套大气再分析资料和观测资料发现 SRP 的位相和振幅存在显著的年代际振荡, 并可以解释总方差的约 50%。在年代际时间尺度上, SRP 具有与年际时间尺度相似的空间结构,但其经向 空间尺度更大,影响范围更广,这可能与年代际时间尺度上 SRP 与极锋急流遥相关的耦合有关(见第 5 节)。 SRP 分别在 20 世纪 70 年代初和 90 年代末经历了两次显著的年代际转折,这对欧亚大陆的夏季气候有重要 影响(Wang et al. 2017)。例如, SRP 的位相在 90 年代末发生了年代际的由正转负,这可以解释欧亚大陆中 纬度地区约 60%的地表气温变化和 40%的降水变化(Wang et al. 2017),也是导致同时期欧亚大陆夏季地表 非均匀增暖(Hong et al. 2017)和东北亚夏季降水减少(Piao et al. 2017)的直接原因。SRP 年代际变化的产 生原因尚不十分明确,一个可能的驱动因子是大西洋多年代际振荡,这得到了一些统计分析和数值试验的 支持(Si and Ding 2016; Hong et al. 2017; Wang et al. 2017; Stephen et al. 2019)。除均值外,SRP 的空间结 构在 70 年代末和 90 年代末也发生了显著的年代际改变,并对与其相联系的区域气候产生了相应的影响 (Wang et al. 2012; Lin et al. 2020)。在上述年代际变化中,SRP 均没有表现出明显的长期 趋势,这说明 SRP 的年代际变化主要来源于气候系统内部变率,而受气候系统外强迫的影响比较小。

4 极锋急流上的大气遥相关

尽管极锋急流的波导效应直到近年来才被学界所注意(见第2节),但关于极锋急流遥相关的工作却可 追溯至20世纪关于鄂霍茨克海高压的研究。鄂霍茨克海高压是位于鄂霍茨克海上方的准定常浅薄高压系统, 当它偏强时会将北方相对寒冷的空气向南输送,造成日本及周边地区的夏季气温偏低,进而影响水稻灌浆、 导致水稻减产(Ninomiya and Mizuno 1985a, 1985b; Kodama 1997),因此其变化受到日本学者的关注。Wang (1992)指出鄂霍茨克海高压的变化伴有横跨欧亚大陆北部的准静止行星尺度波列; Nakamura and Fuamachi (2004)证实了上述结论,并进一步指出该波列对鄂霍茨克海高压的影响存在显著的季节性变化;之后的 许多学者在研究欧亚大陆不同地区的夏季气候异常时也注意到欧亚大陆北部的准静止行星波列结构 (Fukutomi et al. 2003; Wakabayashi and Kawamura 2004; Iwasaki et al. 2006; Iwao and Takahashi 2008; Sun et al. 2015; Li and Ruan 2018; Li et al. 2019)。上述研究从区域气候异常的角度出发给出了欧亚大陆北部准 定常波列存在的证据,但均未注意到极锋急流在这类波列形成中的作用,也缺乏从大气波动或遥相关角度 对波列结构和机制的认识。为此,Xu et al. (2019)从极锋急流波导和大气波动的角度出发,首次定义了夏 季沿极锋急流传播的主要遥相关型。

Xu et al. (2019)利用 250hPa 经向风在欧亚大陆北部 Rossby 波活动最强的区域(图 2c 中的蓝框)内的 EOF1 来定义欧亚大陆上主导的极锋急流遥相关。该遥相关主要由位于英国、波罗的海、西西伯利亚和贝加 尔湖的四个活动中心组成,水平结构表现为沿极锋急流传播的准静止 Rossby 波(图 4a, 4b),垂直结构表 现为在对流层中下层略微西倾的准相当正压分布(图 4c, 4d)。与 SRP 类似,该遥相关的异常高、低压中 心分别对应对流层内的暖、冷温度异常,且温度场略微落后于高度场。由于该遥相关型的第一个和最后一 个中心分别位于英国和贝加尔湖附近,且呈现东西狭长的结构,因此 Xu et al. (2019)将其命名为"英国-贝加尔湖走廊"(British-Baikal Corridor, BBC)遥相关。

BBC 遥相关对欧亚大陆的夏季降水有重要影响。当它处于正位相(图 4a)时,会引起位于其活动中心附近的欧洲东部和东北亚降水偏多、中西伯利亚降水偏少(图 4e),并常伴有极端降水的发生(Hong et al. 2022);同时,BBC 遥相关还会引起长江流域、韩国和日本南部的降水变化(图 4e)。例如,在热带系统和热带外系统的共同作用下,2020年东亚夏季风降水量突破历史记录,并伴有史无前例的"暴力梅"事件(Liu et al. 2020; Takaya et al. 2020)。其中热带系统主要表现为海温异常所激发的西北太平洋低层反气旋性环流响应,这使得西太副高增强、充足的暖湿空气向北输送(Zhou et al. 2020);热带外系统则表现为活跃的BBC

遥相关结构,这有利于引导中高纬地区的干冷空气南下,与向北输送的暖湿空气在长江流域对峙,从而为 "暴力梅"的发生提供了有利的大尺度环流条件 (Park et al. 2021)。

BBC 遥相关对欧亚大陆北部,特别是东欧、中西伯利亚和东西伯利亚地区的平均气温和极端温度事件 也有显著影响(图 4f; X. Li et al. 2021)。例如,2019年6月欧洲中部发生了一次破纪录的热浪,无论是在 月尺度还是日尺度上,BBC 遥相关都是此次事件的主要成因(Xu et al. 2020b)。其作用一方面体现在BBC 遥相关会引发向北的暖平流和下沉增温,从而通过动力作用使地表温度升高,另一方面体现在BBC 遥相关 会减少云量和增加入射短波辐射,以及减小土壤湿度和增加地表感热通量,从而通过热力作用使得地表温 度进一步升高,其中动力作用对地表气温异常的贡献超过 60%(Xu et al. 2021b)。BBC 遥相关对极端温度 的影响不仅体现在极端高温上,而且体现在极端低温上(Xu et al. 2021a)。当包括BBC 遥相关在内的极锋 急流遥相关活动偏强时,欧亚大陆北部特别是欧洲、俄罗斯西部和东北亚地区更容易出现极端温度事件; 而当遥相关活动偏弱时,欧亚大陆北部则不容易出现极端温度事件。这种关系在再分析资料和不同的气候 模式中均非常显著(Xu et al. 2021a)。

与 SRP 不能单纯地以 Rossby 波传播来解释类似,BBC 遥相关在年际尺度上的维持也需要大气不稳定 过程的贡献(Xu et al. 2019)。动力学和能量学分析表明,BBC 遥相关在传播过程中会与背景流发生相互作 用,并主要通过斜压过程从背景流中汲取有效位能来维持自身结构(Xu et al. 2019, Li et al. 2020)。正压过 程引起的动能转换虽然对维持 BBC 遥相关的空间结构贡献很小,但却是决定 BBC 遥相关空间位置的一个 重要原因,天气尺度瞬变扰动的非线性反馈作用对 BBC 遥相关的维持也有正贡献,但湿动力过程对 BBC 遥相关的维持起耗散作用(Xu et al 2019; 2022a)。这些结果表明,BBC 遥相关本质上是极锋急流内的一个 干动力模态,它可以通过干动力过程从背景流中汲取能量以维持自身的结构。

BBC 遥相关主要由北大西洋急流出口区基本气流-低频扰动-高频扰动间的多尺度相互作用过程激发(Xu et al. 2019)。当北大西洋急流出口区存在一个初始低频扰动时,强的背景风场切变会通过正压不稳定使该扰动迅速增强,增强的低频扰动会引起北大西洋风暴轴的经向位移,进而通过风暴轴内天气尺度斜压波的动力强迫作用进一步加强低频扰动,从而激发 BBC 遥相关(Xu et al. 2019)。这种多尺度相互作用依赖于北大西洋急流轴和北大西洋风暴轴的空间位置,因此它与前述的正压不稳定过程一起,共同决定了 BBC 遥相关得以在当前观测中的位置上形成和发展。在年际变化尺度上,海温、海冰和积雪等大气外强迫信号与 BBC 遥相关之间的联系比较弱,但前期冬季北大西洋涛动的正位相有利于夏季 BBC 遥相关呈正位相,这在一定程度上可以作为 BBC 遥相关年际变化的可预报性来源(Xu et al. 2019)。这种大气环流跨季节联系的机制可能与北大西洋环流与北极海冰间的相互作用有关(Ogi et al. 2004),但具体物理过程仍需进一步研究。

在季节内时间尺度上,BBC 遥相关的三维结构与年际尺度基本一致,从其演变可以清楚地看到沿北大 西洋急流东传的初始扰动、初始扰动在北大西洋急流轴出口区的发展、扰动能量向下游欧亚大陆频散形成 BBC 遥相关、以及 BBC 遥相关的衰减等整个过程(Xu et al. 2020a)。这种发展和衰亡主要受非线性过程控 制:天气尺度瞬变波在增长阶段通过正压过程引起 BBC 遥相关增强,而在衰亡阶段则通过斜压过程引起 BBC 遥相关的减弱,这与理想斜压波的生命过程(Simmons and Hoskins 1978)恰好相反。由此可以推测,BBC 遥相关的季节内演变受到低频 BBC 遥相关和高频瞬变波动之间能量循环过程的驱动,反映了非线性过程主 导的波动能量在不同尺度扰动间的能量串级过程。与 SRP 相比,BBC 遥相关的季节内演变表现出更强的非 线性特征,这可能与中高纬地区背景态风速小所导致的非线性效应增强有关(Luo et al. 2019)。线性过程特 别是从基本流中汲取有效位能的斜压过程在 BBC 遥相关的整个生命期中都有利于 BBC 遥相关增强,因此 对 BBC 遥相关的发展起推动作用,而对其衰减起抑制作用。可见,与年际尺度上 BBC 遥相关主要受控于 线性的波流相互作用过程不同,BBC 遥相关在季节内尺度的演变主要受到非线性波波相互作用过程的控制。 以上 BBC 遥相关的季节内演变周期约为 2 周,对应约 4-5 天的指数衰减时间,短于北大西洋涛动和太平洋-北美遥相关,这是由于北大西洋涛动和太平洋-北美遥相关的时间尺度主要依赖于 Rossby 波的破碎和混合 (Feldstein 2002; Benedict et al. 2004; Franzke and Feldstein 2004),而 BBC 遥相关的时间尺度更多地依赖 于波包的群速度和 BBC 遥相关与高频瞬变波动之间的非线性相互作用(Xu et al. 2020a)。

5 副热带急流遥相关与极锋急流遥相关的耦合

副热带急流和极锋急流是两支独立的急流波导,但它们所捕获的大气遥相关在季节内尺度上(Iwao and Takahashi 2008)或有些年份和月份中却可以同时出现,并在欧亚大陆上空呈现双波列结构(图 1a-1c),这暗示着极锋急流遥相关和副热带急流遥相关之间可能存在某种耦合。Xu et al. (2022a)对这种耦合关系的稳定性和物理机制进行了深入分析,发现极锋急流上的第二主模态"英国-鄂霍茨克海走廊"(British-Okhotsk Corridor, BOC)遥相关与副热带急流上的第一主模态 SRP 之间存在随时间变化的耦合关系。BOC 遥相关与BBC 遥相关类似,也是夏季沿极锋急流传播的大气遥相关,它的第一和最后一个活动中心分别位于英国和鄂霍茨克海附近(图 5b),并对欧亚大陆北部的气温和降水有显著影响(图 5d, 5f)。虽然 BOC 遥相关和 BBC 遥相关的空间结构类似,但研究表明无论是在年际时间尺度还是季节内时间尺度上两者之间的相关性均很小,因此可以视为极锋急流上相互独立的准静止大气 Rossby 波(Xu et al. 2022a)。BOC 遥相关会与 SRP 耦合在一起,在欧亚大陆上呈现南北两支的双波列结构(图 5a)。此时,与 BOC 遥相关相联系的显著气候异常不再局限于欧亚大陆北部,而是向南扩展到中纬度和副热带地区(图 5c, 5e),欧亚地区夏季气温(Xu et al. 2022a)和极端高温日数(Deng et al. 2018)年际变率的主模态也会具有更大的经向尺度。这些结果均表明,BOC 遥相关和 SRP 之间的耦合会增强欧亚地区气候异常在中纬度和高纬度之间的联系。

在线性理论框架下,基本气流的结构和遥相关初始扰源的位置均有可能影响 BOC 遥相关和 SRP 之间的 耦合。一系列数值试验表明(Xu et al. 2022a), BOC 遥相关初始扰源的位置对上述耦合起关键作用,而基本 气流差异所起的作用非常小:当初始扰源偏向北大西洋东部时会产生强耦合的双波列结构,而初始扰源偏 向北大西洋西部时会大大削弱沿副热带急流传播的波列结构,从而形成主要沿极锋急流传播的 BOC 遥相关。 进一步分析表明,初始扰源的位置受北大西洋三极型海温异常和北大西洋急流出口区波流相互作用过程的 影响。

6 气候变化对急流遥相关的影响

当前,人类活动引起的气候变化已经引起大气、陆地、海洋等气候系统成员的广泛变化,并使极端天 气气候事件更加频繁和严重 (IPCC 2021), 其中许多中高纬地区的极端天气事件均伴有急流遥相关的活动 (Screen et al. 2014; Xu et al. 2021a),这引发了关于气候变化与急流遥相关活动之间因果关系的猜测。目前 关于气候变化是否及如何影响急流遥相关主要有两方面观点,第一类观点认为气候变化会改变大气环流的 基本态,这将使大气行星波的共振现象更容易出现,从而有利于急流遥相关振幅的增加。该观点的基础来 源于 Petoukhov et al. (2013) 提出的"准共振放大"理论,即大气内部纬向波数为 6-8 的自由波可以和受到 稳定地形或热力强迫激发的定常波之间发生共振,从而使沿急流波导传播的准静止行星波在短时间内迅速 增长并形成高振幅的急流遥相关。由于在全球变暖的气候背景条件下一些被急流波导所捕获的自由波会因 为背景环流结构的变化而趋向于准静止,因此准共振放大现象将在未来更容易发生(Coumou et al. 2018)。 然而,虽然已有不少学者用该理论来解释近年来不断增多的极端天气事件(Coumou et al. 2014; Kornhuber et al. 2017a, 2017b, 2019) 和未来急流遥相关的可能变化(Mann et al. 2017, 2018; Coumou et al. 2018), 但 仍有很多学者由于以下三方面原因对准共振放大理论的合理性和真实性存在疑虑:第一,准共振放大理论 目前仍停留在对数学方程的简化推导阶段,尚未在理想的动力机制模型中得到系统的验证(Teng and Branstator 2019); 第二, 准共振放大理论强调次季节尺度基本气流对波动振幅的影响, 但在次季节尺度上区 分基本气流和扰动十分困难,尤其是在波动振幅很大的情形(Wirth and Polster 2021); 第三,准共振放大理 论本质上仍是基于线性假设下的物理机理,因此并不能解释波动快速增长中的关键非线性过程(Schubert et al. 2011; Xu et al. 2020a).

关于气候变化是否及如何影响急流遥相关的第二类观点认为,气候变化会通过改变和增强大气中的非绝热加热源从而有利于急流遥相关振幅的增加,这一观点主要来源于 Teng et al. (2019)提出的土壤湿度异

常对急流遥相关的激发作用。基于地球系统模式的数值试验表明,对流层大气对北美地区土壤湿度变化的 响应非常类似于急流遥相关的波列活动(Teng et al. 2019),而随着未来地表温度的升高以及相伴随的潜在蒸 散发的增强,与土壤湿度异常相关的陆气相互作用过程会在许多地区增强(Teng et al. 2016),因此急流遥相 关的活动很可能会在未来地表非绝热加热增强的情况下增加(Teng and Branstator 2019)。该观点具有较合理 的物理机制并得到数值试验的支持,但仍需进一步验证,这是因为急流遥相关对土壤湿度异常的响应依赖 于所使用的模式和试验设计。例如,Koster et al. (2016)在另一个大气环流模式中使用了一种更符合实际物 理过程的方式改变局地土壤湿度,但模式并未出现 Teng et al. (2019)中的急流遥相关响应。此外,当前模 式中的许多陆气相互作用过程都依赖于经验性的参数化方案,这也会导致模式在预估未来的地表土壤湿度 和相关地表非绝热加热变化时存在很大的不确定性。值得注意的是,无论是基本气流的变化还是非绝热加 热强迫的改变,急流遥相关的振幅在最近几十年中并没有表现出显著的趋势(Screen and Simmonds 2013), 且模式预估的副热带急流附近的年际扰动和次季节扰动振幅在未来都会随着经向温度梯度的减小而减弱 (Teng et al. 2016)。因此,关于气候变化如何影响急流遥相关的变化及其机理目前仍未有明确统一的认识, 需要在未来进一步深入研究。

7 总结与展望

本文聚焦夏季急流遥相关活动最为活跃的欧亚大陆地区,简要总结了近年来关于夏季急流遥相关的研究进展。研究表明,副热带急流和极锋急流是两支独立的可供大气Rossby波传播的波导,其中前者的波导效应主要由强风速切变形成,而后者的波导效应主要由强层结变化形成。"丝绸之路"遥相关SRP 是沿副热带急流波导传播的主要大气遥相关,它主要在线性的大气内动力过程和大气外强迫变化的共同作用下得以维持;受气候系统内部变率的控制,SRP 呈现显著的年代际变化,并影响欧亚大陆中纬度地区的平均和极端降水、平均气温和极端热浪。"英国-贝加尔湖走廊"(BBC)遥相关和"英国-鄂霍茨克海走廊"(BOC)遥相关是沿极锋急流波导传播的主要大气遥相关,它们主要受非线性大气内动力过程的控制,其中北大西洋急流出口区的基本气流-低频波-高频波多尺度相互作用是其主要激发机制,而遥相关自身与基本气流间的斜压能量转换过程是其主要维持机制。BBC和BOC 遥相关可以显著影响欧亚大陆高纬度地区的平均和极端降水事件、平均气温和极端温度事件,且BOC 遥相关与SRP 之间存在具有年代际变化的耦合关系,其耦合与否对欧亚大陆中高纬度地区的夏季气候和变率有重要影响,北大西洋上空遥相关初始扰源的位置是决定二者耦合与否的关键。急流遥相关的活动在气候变化背景下可能会加强,气候变化会对大气环流基本态的改变及其引起的大气行星波共振的变化以及气候变化对非绝热加热源的改变是产生这种变化的可能原因。

急流遥相关作为影响中高纬夏季气候的重要环流系统,在中高纬地区夏季天气预报、气候预测乃至气候预估方面发挥着关键作用。尽管目前对急流遥相关的变化特征、机制及其气候影响已有较为系统的认识,但仍存在需要进一步研究的科学问题,从而限制了相应时间尺度上的预测和预估能力。在次季节和年际时间尺度上,需要深入了解大气内动力和大气外强迫对急流遥相关的相对重要性,提升数值模式对其中关键物理过程的模拟能力。在年代际时间尺度上,需要进一步研究 AMO 等大气外强迫影响急流遥相关的物理过程,提升数值模式对这些大气外强迫以及相关气候系统响应的模拟能力;此外,还需要明确为什么决定副热带急流遥相关和极锋急流遥相关之间是否耦合的初始扰源会存在年代际移动,以及数值模式是否能够准确预测急流遥相关之间的耦合。在百年时间尺度上,需要减小数值模式预估基本流和大气外强迫对全球变暧响应的不确定性,并研究急流遥相关在相应基本流和大气外强迫影响下的变化机理。回答这些问题将有助于加深对不同时间尺度上急流遥相关变化特征和机理的认识,从而为提升中高纬地区夏季的天气预报、气候预测乃至气候预估提供理论基础。此外,除了具有准定常波性质的急流遥相关外,欧亚大陆中高纬地区还存在移动性西传的周期约 30-50 天的大气扰动,此类扰动可对乌拉尔山附近的温度产生影响,对此未来也值得进一步研究。

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图 1 250hPa 经向风(阴影)的月平均异常在(a) 2010 年 7 月,(b) 2019 年 6 月和(c) 2020 年 6 月的空间分布(单位: m/s)。(d)-(f)同(a)-(c),但为每日最高地表温度(Tmax,阴影)的月平均异常(单位: °C)。(a)-(c)中的黑色粗虚线代表 250hPa 纬向风的气候态分布超过 15 m/s 的 地区。(d)-(f)中的打点区域表示极端高温事件在该月的发生频率超过气候态 8 倍的地区。引自 Xu et al. (2021a)。

Fig.1 The monthly mean 250-hPa meridional wind (shading) anomalies in (a) July 2010, (b) June 2019, and (c) June 2020, respectively. (d)-(f) are the same as (a)-(c), but for the monthly mean of daily maximum surface temperature (Tmax, shading). Contours in (a)-(c) indicate the elimatological 250-hPa zonal wind speed that exceeds 15 m/s. Stipplings in (d)-(f) indicate the frequency of extreme high temperature events is eight times more than elimatology in the specific month. Adapted from Xu et al. (2021a).





图 2 (a)标准化后的夏季 250-hPa 绝对涡度经向梯度的气候态分布。(b)标准化后的夏季 250-hPa 位涡经向梯度的气候态分布。(c)标准化后的夏季 250-hPa 经向风的年际标准差分布。图 (a) - (c)中的标准化是指在每个格点值都除以了每个格点经度上从 0°到 90°N 的面积加权值。(c)中的蓝色 框 (20°W-150°E, 50°N-80°N)和红色框 (30°E-130°E, 20°N-60°N)分别代表极锋急流遥相关和副热带急流遥相关的定义区域。引自 Xu et al. (2022a)并重绘。

Fig.2 (a) The standardlized meridional gradient of the climatological-mean summer 250-hPa absolute vorticity. (b) The standardlized meridional gradient of the climatological-mean summer 250-hPa potential vorticity. (c) The standardlized temporal standard deviation of the summer 250-hPa meridional wind. The standardlization in (a) - (c) means the value in each grid point has been divided by the area-weighted mean from 0°-90°N at its longitude. The blue and red boxes in (c) represent the region (20°W-150°E, 50°N-80°N) and the region (30°E-130°E, 20°N-60°N), where the waveguide teleconnections along the polar front jet and subtropical jet are defined, respectively. Adapted from Xu et al. (2022a)

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图 3 (a) SRP 指数回归的夏季平均的 200-hPa 位势高度场 (黑色等值线,等值线间隔: 10gpm) 以及相应的水平波活动通量 (箭头,单位: m² s⁻²)。 紫色等值线代表 200-hPa 夏季平均的纬向风速为 20m s⁻¹的等值线。(c) SRP 指数回归的夏季平均的位势高度场 (黑色等值线,等值线间隔: 10gpm)、 温度场 (蓝色等值线,等值线间隔: 0.3°C) 和波作用通量 (矢量,单位: m² s⁻²) 在 30°N 至 50°N 内平均的垂直剖面图。其中 (c) 中波作用通量的水 平分量为波作用通量纬向分量和经向分量的平方根。为了更好的展示效果,波作用通量的垂直分量放大了 120 倍。虚线代表负值,零线没有画出。图 (a) - (c) 中浅色阴影和深色阴影分别代表双尾 Student's t 检验通过了 95%和 99%的信度水平。图 (b) - (d) 与 (a) - (c) 相同,但是为相对涡度 场的空间分布。(e) SRP 指数回归的夏季平均的降水场 (阴影间隔: 2 mm month⁻¹)。(f) SRP 指数回归的夏季平均的地表温度 (阴影间隔: 0.2°C)。 图 (e) - (f) 中灰色和黑色粗点分别代表双尾 Student's t 检验通过了 95%和 99%的信度水平。引自 Wang et al. (2017) 并重绘。

Fig.3 (a) Summer mean 200 hPa geopotential height anomalies [black contour, contour interval (CI)=10 gpm] and the horizontal component of the wave activity flux (arrow, unit: $m^2 s^{-2}$) associated with the SRP index, obtained via a linear regression overlaid by the climatology of the 200 hPa summer mean zonal wind, as indicated by the 20 m s⁻¹ purple contour. (c) The regression of the 30°-50°N averaged summer mean geopotential height (black contour, CI=5 gpm) and air temperature (purple contour, CI=0.2 °C) onto the SRP index. (b) and (d) are the same as (a) and (c), but geopotential height is replaced with relative vorticity [black contour, CI=4 ($\pm 2, \pm 6, \pm 10,...$) × 10⁻⁶ s⁻¹ for (b) and CI=2 ($\pm 1, \pm 3, \pm 5,...$) × 10⁻⁶ s⁻¹ for (d)]. Regression of the summer mean (e) precipitation (SI=2 mm month⁻¹), and (f) surface air temperature (SI=0.2 °C) onto the SRP index. The light and dark shading in (a)-(d) and the gray and black dots in (e)-(f) indicate the 95% and 99% confidence levels based on two-tailed Student's *t*-test, respectively. Adapted from Wang et al. (2017).

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图 4 (a) BBC 指数回归的夏季平均的 250-hPa 位势高度场 (黑色等值线,等值线间隔: 10gpm)以及相应的水平波活动通量 (箭头,单位: m² s⁻²)。 紫色等值线代表 250-hPa 夏季平均的纬向风速为 20m s⁻¹ 的等值线。(c) BBC 指数回归的夏季平均的位势高度场 (黑色等值线,等值线间隔: 10gpm)、 温度场 (蓝色等值线,等值线间隔: 0.3°C)和波作用通量 (矢量,单位: m² s⁻²)在 55°N 至 65°N 内平均的垂直剖面图。其中 (c)中波作用通量的水 平分量为波作用通量纬向分量和经向分量的平方根。为了更好的展示效果,波作用通量的垂直分量放大了 120 倍。虚线代表负值,零线没有画出。图 (a) - (c)中浅色阴影和深色阴影分别代表双尾 Student's t 检验通过了 95%和 99%的信度水平。图 (b) - (d)与 (a) - (c)相同,但是为相对涡度 场的空间分布。(e) BBC 指数回归的夏季平均的降水场 (阴影间隔: 2 mm month⁻¹)。(f) BBC 指数回归的夏季平均的地表温度 (阴影间隔: 0.2°C)。 图 (e) - (f)中灰色和黑色粗点分别代表双尾 Student's t 检验通过了 95%和 99%的信度水平。引自 Xu et al. (2019)并重绘。

Fig.4 (a) Summer mean 250 hPa geopotential height anomalies [black contour, contour interval (CI)=10 gpm] and the horizontal component of the wave activity flux (arrow, unit: $m^2 s^{-2}$) associated with the BBC index, obtained via a linear regression overlaid by the climatology of the 250 hPa summer mean zonal wind, as indicated by the 20 m s⁻¹ purple contour. (c) The regression of the 55°-65°N averaged summer mean geopotential height (black contour, CI=5 gpm) and air temperature (purple contour, CI=0.2 °C) onto the BBC index. (b) and (d) are the same as (a) and (c), but geopotential height is replaced with relative vorticity [black contour, CI=4 ($\pm 2, \pm 6, \pm 10,...$) × 10⁻⁶ s⁻¹ for (b) and CI=2 ($\pm 1, \pm 3, \pm 5,...$) × 10⁻⁶ s⁻¹ for (d)]. Regression of the gray and black dots in (e)-(f) indicate the 95% and 99% confidence levels based on two-tailed Student's *t*-test, respectively. Adapted from Xu et al. (2019).

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图 5 (a) 高相关时段内 BOC 指数回归的夏季平均的 250-hPa 位势高度场 (黑色等值线,等值线间隔: 10gpm) 以及相应的水平波活动通量 (箭头,单 位: m² s⁻²)。紫色等值线代表 250-hPa 夏季平均的纬向风速为 18m s⁻¹ 的等值线。(c) 高相关时段内 BOC 指数回归的夏季平均的降水场 (填色,填色间 隔: 2 mm/月)。(e) 高相关时段内 BOC 指数回归的夏季平均的地表气温场 (填色,填色间隔: 0.2 °C)。(b)、(d)、(f) 同 (a)、(c)、(e),但为低相 关时段。(a)-(b)中的浅色和深色阴影以及 (c)-(f)中的灰色和黑色点表示基于 Student's *t* 检验通过了 90%和 95%的信度水平。引自 Xu et al. (2022a)。 Fig.5 Summer mean Z250 anomalies (black contour, CI = 5gpm) and the horizontal component of the wave activity flux (arrow, unit: m² s⁻²) associated with the BOC index during (a) the HIGH- and (b) the LOW-epochs, obtained via a linear regression overlaid by the climatology of the U250, as indicated by the 18m s⁻¹ purple contour. (c) and (d) are same as (a) and (b), but for the precipitation [shading, shading interval (SI) = 2 mm month ⁻¹]. (e) and (f) are the same as (a) and (b), but for the surface air temperature (shading, SI = 0.2 °C). The light and dark shading in (a)–(b) and the gray and black dots in (c)-(f) indicate the 90% and 95% confidence levels based on the two-tailed Student's *t*-test, respectively. Adapted from Xu et al. (2022a).